1 2	The impact of shipping, agricultural, and urban emissions on single particle chemistry observed aboard the R/V Atlantis during CalNex
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4 5	Cassandra J. Gaston, <sup>1‡</sup> Patricia K. Quinn, <sup>2</sup> Timothy S. Bates, <sup>2</sup> Jessica B. Gilman, <sup>3</sup> Daniel M. Bon, <sup>3†</sup> William C. Kuster, <sup>3</sup> and Kimberly A. Prather <sup>1,4*</sup>
6	
7	<sup>1</sup> Scripps Institution of Oceanography, University of California, San Diego, La Jolla, CA 92093 USA
8	
9	<sup>2</sup> NOAA Pacific Marine Environmental Laboratory, Seattle, WA 98115 USA
10	
11 12	<sup>3</sup> NOAA Earth System Research Laboratory, Boulder, CO 80305 USA and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, CO 80309
13	
14 15	<sup>4</sup> Department of Chemistry and Biochemistry, University of California, San Diego, La Jolla, CA 92093- 0314 USA
16	
17 18	<sup>‡</sup> Current Address: Department of Atmospheric Sciences, University of Washington, Seattle, WA 98195 USA
19	
20 21	<sup>†</sup> Current Address: Planning and Policy Program, Air Pollution Control Division, Colorado Department of Public Health and Environment, Denver, CO 80246 USA
22	
23	*Corresponding author: kprather@ucsd.edu, 858-822-5312. Fax: 858-534-7042.
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25 Abstract:

The CalNex field campaign was undertaken to obtain a better understanding of the 26 regional impacts of different pollution sources in California. As part of this study, real-time 27 shipboard measurements were made of the size-resolved single-particle mixing-state of sub- and 28 29 super-micron particles (0.2-3.0 um aerodynamic diameter) along the California coast where major differences were noted between Southern and Northern California. In Southern California, 30 particles containing soot made up the largest fraction of submicron particles (~38% on average 31 32 and up to ~89% by number), whereas organic carbon (OC) particles comprised the largest fraction of submicron number concentrations ( $\sim 29\%$  on average and up to  $\sim 78\%$  by number) in 33 Northern California including the Sacramento area. The mixing-state of these carbonaceous 34 particle types varied during the cruise with sulfate being more prevalent on soot-containing 35 particles in Southern California due to the influence of fresh shipping and port emissions in 36 addition to contributions from marine biogenic emissions. Contributions from secondary organic 37 aerosol species, including amines, and nitrate were more prevalent in Northern California, as 38 well as during time periods impacted by agricultural emissions (e.g., from the inland Riverside 39 40 and Central Valley regions). These regional differences and changes in the mixing-state and sources of particles have implications for heterogeneous reactivity, water uptake, and cloud 41 nucleating abilities for aerosols in California. 42

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**Index Terms:** 0300, 0305, 0345, 0365

46 **1. Introduction** 

Atmospheric aerosols contribute to air pollution, have adverse effects on human health, 47 and perturb the Earth's radiative balance [Poschl, 2005]; both aerosol size and composition play 48 key yet uncertain roles in these effects. One aim of the Research at the Nexus of Air Quality and 49 50 Climate Change (CalNex) field campaign was to elucidate the links between aerosols, air pollution, and climate in order to guide policies regarding emission regulations in California. 51 Diverse particle sources impact California including ships, vehicle exhaust, oil refineries, meat 52 53 cooking, and animal husbandry and agricultural emissions. The Ports of Los Angeles (LA) and Long Beach (LB) are the busiest container ports in the United States, contributing high levels of 54 ship and port emissions (e.g., diesel truck emissions, oil refinery emissions) to Southern 55 California. Locations in inland Southern California (e.g., Riverside) and the Sacramento and San 56 Joaquin Valley areas in Northern California are impacted by dairy farm and agricultural 57 emissions leading to high mass concentrations of particulate matter that contain large fractions of 58 secondary species [Chen et al., 2007; Chow et al., 2006a; Docherty et al., 2008; Grover et al., 59 2008; Hughes et al., 2000; Hughes et al., 2002; Magliano et al., 1999; Pastor et al., 2003; Oin et 60 al., 2012; Sorooshian et al., 2008]. Further, biogenic emissions from forested regions such as 61 the Sierra Nevada foothills and from marine biological activity along the California coast also 62 contribute to the aerosol burden in California [Creamean et al., 2011; Gaston et al., 2010]. 63 64 Assessing how these regional differences in particle sources and secondary reactions in California impact the physicochemical properties of aerosols represents a key step in fulfilling 65 the goals of the CalNex field campaign. 66

67 Single-particle mass spectrometry is well-suited for providing the high temporal
 68 resolution and mass spectral fingerprints necessary for distinguishing diverse particle sources as

69 well as assessing the impact of atmospheric processing (e.g., gas-to-particle partitioning, heterogeneous reactions, etc.) on particle size and chemistry [Pratt and Prather, 2011; Sullivan 70 and Prather, 2005]. Aerosol time-of-flight mass spectrometry (ATOFMS), a single-particle 71 technique, has been successfully used for the source apportionment of particulate matter in 72 California for almost fifteen years [Ault et al., 2010; Creamean et al., 2011; Gaston et al., 2010; 73 Hughes et al., 2000; Moffet et al., 2008b; Noble and Prather, 1996; Pastor et al., 2003; Pratt et 74 al., 2009; Pratt and Prather, 2009; Oin et al., 2012; Toner et al., 2008; Whiteaker et al., 2002]. 75 Herein we present real-time measurements of single-particle composition and size using 76 ATOFMS during the CalNex campaign onboard the R/V Atlantis sampling platform. 77 Atmospheric measurements were made off the California coast, targeting specific sources, 78 including the Ports of LA and LB, continental outflow from the LA Basin, and emissions from 79 80 Northern California including the inland, Sacramento area. This paper describes the significant differences observed in single-particle mixing-state and sources between these regions. The 81 implications of these findings are discussed. 82

- 83 2. Experimental
- 84 2.1 CalNex

Ambient aerosol and gas-phase measurements were made onboard the R/V Atlantis from May 14-June 8, 2010 as part of the CalNex 2010 field campaign ([*Ryerson et al.*, 2013]; <u>http://www.esrl.noaa.gov/csd/calnex/; http://saga.pmel.noaa.gov/data</u>). The ship traveled from San Diego to Sacramento and then back to the Port of San Francisco where the study ended. The cruise track is shown in Figure 1. The air in the aerosol sampling manifold was conditioned to  $60 \pm 5\%$  relative humidity (RH) using a heated inlet. Insulated conductive tubing extending

91 from the heated inlet provided air to the instruments at flows of  $\sim 30$  lpm [Bates et al., 2004]. Meteorological parameters, aerosol properties, and gas-phase constituents such as radon, SO<sub>2</sub>, 92 O<sub>3</sub>, etc. were measured onboard the ship [*Quinn and Bates*, 2005]. Hydrocarbons, including 93 toluene and benzene, were quantified using gas chromatography equipped with a flame 94 ionization detector (GC-FID) [Bon et al., 2011]. The ratio of toluene to benzene can be used 95 as an indicator of photo-chemical age [Gelencser et al., 1997; Roberts et al., 1984]. Air masses 96 containing fresh urban emissions that have undergone minimal photochemical processing 97 typically have toluene/benzene ratios  $\geq 2$ , lower values are indicative of more photo-chemically 98 processed air masses as the more reactive toluene is preferentially removed by reactions with the 99 hydroxyl radical. Clean marine air masses are typically well-aged and are expected to have 100 toluene/benzene ratios <<1. Ship exhaust from the R/V Atlantis was filtered out by eliminating 101 time periods when the relative wind direction was aft of +100° (starboard) and -120° (port) off 102 the bow. All data presented here were filtered using this method. Data are presented in 103 coordinated universal time (UTC), which is 7 hours ahead of local time (PDT), as day of year 104 105 (DOY). For example, a date and time of 12:00 on January 1 presented in UTC would correspond to DOY 1.5. 106

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#### 2.2 Aerosol Measurements: ATOFMS

108 The size-resolved chemical composition of individual aerosol particles from 0.2-3.0 um 109 aerodynamic diameter was measured in real-time using an aerosol time-of-flight mass 110 spectrometer (ATOFMS). The operating principles of the ATOFMS have been described 111 previously [*Gard et al.*, 1997; *Prather et al.*, 1994]. Briefly, particles are sampled through a 112 converging nozzle inlet into a differentially pumped vacuum chamber causing particles to be

accelerated to a size-dependent terminal velocity. Particles next enter the sizing region of the 113 114 instrument consisting of two continuous wave (532 nm) lasers separated by a fixed distance. The time taken to traverse the laser beams is recorded giving the terminal velocity of the particle, 115 116 which is used to calculate the aerodynamic diameter of the particle. The particle velocity is also used to time the firing of a Q-switched Nd:YAG laser (266 nm) operating at ~1.2 mJ laser power 117 that simultaneously desorbs and ionizes compounds from individual particles creating positive 118 and negative ions, which are analyzed in a dual polarity time-of-flight mass spectrometer. 119 Particle detection efficiencies are dependent on the ability of particles to interact with the 266 nm 120 radiation [Dall'Osto et al., 2006; Oin et al., 2006; Wenzel et al., 2003]. Dual-polarity spectra 121 provide complementary information regarding the source (e.g., ships vs. sea salt) and age of the 122 particle (e.g., fresh vs. reacted sea salt) [Guazzotti et al., 2001; Noble and Prather, 1996]. A lack 123 124 of negative ion spectra indicates the presence of tightly bound particle phase water, which suppresses negative ions produced by laser desorption/ionization [Neubauer et al., 1997; 125 Neubauer et al., 1998]. Conditioning particles to 60% RH decreases some of this water, thus, 126 increasing the formation of negative ion spectra. 127

128 The YAADA software toolkit was used to import ion peak lists into MATLAB v 6.5.1 (The MathWorks, Inc.) for processing of ATOFMS data [Allen, 2002]. Single-particle mass 129 spectra were then analyzed using a clustering algorithm (ART-2a), which groups spectra together 130 with similar characteristics into distinct "clusters" [Song et al., 1999]. These "clusters" are then 131 merged into distinct particle types based on the prevalent mass spectral ions and intensities, 132 which are indicative of particle sources and chemistry [Noble and Prather, 1996]. 133 The prevalence of these particle types were averaged into 1-hour time bins and separated into 134 submicron (0.2-1.0 um) and supermicron (1.0-3.0 um) particles. Each ion peak assignment 135

presented in this paper corresponds to the most likely ion produced at a given mass-to-charge (m/z). Particle types described herein are defined by characteristic ion peaks and/or possible sources and do not reflect all of the species present within a particular particle class.

139 **3. Results** 

Temporal trends of single-particle measurements were analyzed to discern differences in 140 particle composition between sources in Southern and Northern California. The temporal 141 variability of the most frequently observed particle types detected by ATOFMS as well as the 142 latitudinal position of the ship are shown in Figures 2a and 2b for sub- and super-micron 143 particles, respectively. Several trends in single-particle mixing-state were identified based on 144 differences in particle source, meteorological conditions, and aging processes. To illustrate this, 145 146 six distinct time periods are identified by colored boxes in Figure 2; the time periods are defined as Riverside Transport (Period 1, boxed in red), Stagnant/Ports Transport (Period 2, boxed in 147 black), Marine/Coastal Transport (Period 3, boxed in cvan), Ports of Los Angeles/Long Beach 148 149 (LA/LB) (Period 4, boxed in green), Inland/Valley Transport (Period 5, boxed in orange), and Bay Area/Sacramento (Period 6, boxed in purple). The following sections provide a detailed 150 comparison and discussion of the gas-phase species and meteorological conditions, prevalent 151 particle types and sources, the mixing-state of carbonaceous particles, and the secondary 152 particulate species present during each time period. 153

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#### **3.1** Characteristics of Each Period

During the Riverside, Stagnant/Ports, and Marine/Coastal Transport periods,
measurements took place along the Southern California coast focusing on emissions from the LA

157 Basin, whereas measurements during the Ports of LA/LB period were entirely performed in the Ports of LA and LB and surrounding shipping lanes. During the Inland/Valley Transport period, 158 measurements extended further north to include the Santa Barbara region; the Bay 159 160 Area/Sacramento period measurements were made in Northern California when the ship remained in the Deep Water Channel/Sacramento region for 3 days. Table 1 shows the 161 corresponding dates and meteorological and gas phase measurements for each time period. 162 Figure 3 shows the 48-hour calculated air mass back trajectories at 500 m height [Draxler and 163 *Rolph*, 2011] for each time period to highlight differences in air mass transport conditions. The 164 Stagnant/Ports and Ports of LA/LB periods were heavily influenced by emissions from the Ports 165 of LA/LB while the Riverside, Stagnant/Ports, Inland/Valley, and Bay Area/Sacramento periods 166 were influenced by agricultural and urban emissions from Riverside and/or the Central Valley. 167 168 The Marine/Coastal Transport period was influenced by oceanic emissions and serves as a background period. 169

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## 3.1.1 Period 1: Riverside Transport

During the Riverside Transport period, HYSPLIT analysis indicates that the sampled air 171 masses were transported from the Riverside and Imperial Valley regions before traversing the 172 port and Santa Monica regions (see Figure 3a). Thus, the Riverside Transport period is expected 173 to be influenced by urban, port, and agricultural emissions. Particulate matter in Riverside 174 typically is well-aged as indicated by high concentrations of secondary species such as nitrate, 175 amines, ammonium, and secondary organics [Hughes et al., 2000; Hughes et al., 2002; Liu et al., 176 2000; Pastor et al., 2003; Pratt et al., 2009; Pratt and Prather, 2009; Qin et al., 2012]. The ratio 177 of toluene to benzene was not available during this period; however, available NO<sub>x</sub>/NO<sub>y</sub> ratios 178

indicated that the air masses during this period were photo-chemically aged. The average radon concentration was high  $(2500 \pm 760 \text{ mBq/m}^3)$  indicating that continentally influenced air masses were sampled. The average temperature was low  $(13.2 \pm 0.6^{\circ}\text{C})$  while RH was high  $(89 \pm 3\%)$ .

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## **3.1.2** Period 2: Stagnant/Ports Transport

Sampled air masses stagnated around the coast near the ports and Santa Monica region 183 during the Stagnant/Ports Transport period. Thus, port emissions (e.g., emissions from vehicles, 184 ships, etc.) contributed more significantly to the aerosol burden during the Stagnant/Ports 185 Transport than the Riverside Transport period. Some air masses also originated from the inland, 186 Central Valley area toward the end of the Stagnant/Ports Transport period potentially carrying 187 agricultural emissions. The lowest average toluene to benzene ratio, 0.4, was observed during 188 189 this period suggesting that air masses sampled during this period were heavily photo-chemically aged. Again, high RH (91  $\pm$  3% on average) and low temperatures (13.4  $\pm$  1.1°C) were 190 observed. 191

### 192 **3.1.3** Period 3: Marine/Coastal Transport

During the Marine/Coastal Transport period, most air masses followed a northern coastal/oceanic trajectory along the California coast. Wind speeds reached 14.5 m/s, which enhanced the production of fresh sea spray particles from bursting bubbles generated by breaking waves [*Blanchard and Woodcock*, 1957; *Monahan et al.*, 1983; *O'Dowd and De Leeuw*, 2007]. Thus, the Marine/Coastal Transport period is characterized by ocean-derived aerosol comprised of both fresh sea salt and biogenic organics. This period is not necessarily representative of clean marine conditions based on the relatively high radon concentrations (1550  $\pm$  1080 mBq/m<sup>3</sup>) and high average particle number counts  $(5721 \pm 2680 \text{ cm}^{-3})$  [*Fitzgerald*, 1991; *Hawkins et al.*, 2010; *O'Dowd and De Leeuw*, 2007; *Twohy et al.*, 2005]. Further evidence that the Marine/Coastal Transport period is not entirely a clean marine period also stems from the toluene to benzene ratio, which ranged from 0.030 during clean marine conditions to 3.74 when fresh urban emissions were encountered with an average of 0.85.

## 205 3.1.4 Period 4: Ports of LA/LB

The Ports of LA/LB period was characterized by high  $SO_{2(g)}$  concentrations (3.4 ± 5.2 206 ppbv on average with a maximum of 40 ppbv) compared to the previous periods when average 207  $SO_{2(g)}$  concentrations ranged from 0.06-0.45 ppbv. Further, low wind speeds (2.4 ± 1.2 m/s) and 208 low radon concentrations  $(740 \pm 320 \text{ mBg/m}^3)$  were also encountered. Toluene to benzene ratios 209 were high with an average of 1.89 and a maximum of 6.1 indicating a photo-chemically 210 unprocessed air mass. It is likely that local urban, port, and shipping emissions dominated the 211 particle composition during this period with little influence from other continental or transported 212 213 sources.

# 214 3.1.5 Period 5: Inland/Valley Transport

Figure 3e shows air masses traveling across the Central Valley and the desert before traversing the LA Basin and Santa Barbara regions during the Inland/Valley Transport period. Higher radon and  $O_{3(g)}$  concentrations were measured during this time period compared to other time periods with averages of  $3200 \pm 3000 \text{ mBq/m}^3$  and  $44 \pm 11 \text{ ppbv}$ , respectively. Toluene to benzene ratios had an average of 0.84 indicating that air masses sampled during this period were fairly well-aged. Similar to Riverside, aerosol from the Central Valley is typically characterized by high concentrations of secondary species, namely secondary organic aerosol and ammonium
nitrate due to contributions from dairy farms and other agricultural emissions in addition to urban
emissions from sources such as vehicles [*Chen et al.*, 2007; *Chow et al.*, 2006a; *Chow et al.*,
2006b; *Eatough et al.*, 2008]. Hence, particulate matter observed during this period is expected
to show similar signs of particle aging as the Riverside Transport period.

# 226 3.1.6 Period 6: Bay Area/Sacramento

During the Bay Area/Sacramento period, air masses traveled over the San Francisco Bay area prior to arriving in the Sacramento region, which is within the Central Valley. Diurnal temperature and RH profiles were observed with night and day values ranging from 16 to 30°C and 24 to 93%, respectively. High toluene to benzene ratios were observed with an average of 1.74, suggesting that the sampled air masses were relatively fresh, similar to the Ports of LA/LB period; however, unlike the Ports of LA/LB period, sampling during the Bay Area/Sacramento period occurred within the Central Valley and not within a major port region.

### 234 **3.2 Observed Particle Types**

#### 235 **3.2.1** Submicron Particle Composition

Particle composition varied based on sources and transport conditions (Figures 2 and 4). High number fractions of soot-containing particles (up to ~89% of submicron particles by number with an average of  $\sim 38 \pm 27\%$ ) were observed in Southern California, particularly during the Riverside Transport, Stagnant/Ports Transport, and Ports of LA/LB periods (Periods 1, 2, and 4) when port and urban (e.g., vehicles) emissions were dominant. The main exception to this trend in high soot-containing number fractions in Southern California occurred during the Marine/Coastal Transport period (Period 3) when submicron particle composition was dominated by sea salt particles ( $\sim 63 \pm 17\%$  of submicron particles on average) due to air mass transport conditions. The dominance of soot-containing particles in Southern California is in agreement with aircraft ATOFMS measurements made during CalNex where *Cahill et al.*, [2012] reported high soot-containing number fractions in the Southern California/LA Basin region.

In addition to the soot-containing particle type, particles from residual fuel combustion 247 (e.g., emissions from ships and oil refineries) were also observed namely during the 248 Stagnant/Ports Transport and Ports of LA/LB periods when port and shipping emissions heavily 249 influenced particle composition. This residual fuel combustion particle type represented at most 250  $\sim 25\%$  of submicron particles and is characterized by ion peaks associated with transition metals 251 found in residual fuel oil, notably vanadium ( ${}^{51}V^+$ ,  ${}^{67}VO^+$ ), nickel ( ${}^{58,60}Ni^+$ ), and iron ( ${}^{54,56}Fe^+$ ), in 252 addition to organic carbon, sulfate  $({}^{97}HSO_4)$  and, to a lesser extent, nitrate  $({}^{46}NO_2, {}^{62}NO_3)$  (see 253 Figure 5a for an example mass spectrum); these particles are herein referred to as V-OC particles 254 [Agrawal et al., 2008; Ault et al., 2010; Ault et al., 2009; Corbett and Fischbeck, 1997; Healy et 255 al., 2009; Murphy et al., 2009]. During the Ports of LA/LB period, additional industrial particle 256 257 types were observed, including metals concentrated in the submicron size mode, which most likely represent emissions from incineration [Moffet et al., 2008a] and are the subject of a 258 different manuscript [Weiss-Penzias et al., 2013]. Additionally, organic carbon (OC) 259 represented a much higher fraction of the detected submicron particles during the Ports of 260 LA/LB period ( $\sim 14 \pm 14\%$  of the total submicron particles detected, on average) than previous 261 periods likely due to elevated fossil fuel emissions in the Ports of LA and LB from vehicles, 262 heavy duty diesel trucks and ships, which have been shown to contribute high mass 263

concentrations of organic carbon to the aerosol burden [*Lack et al.*, 2009; *Murphy et al.*, 2009; *Russell et al.*, 2000; *Sodeman et al.*, 2005].

In contrast, soot-containing and V-OC particles made smaller contributions to the number 266 concentration in Southern California during Inland/Valley Transport conditions and in the 267 Northern California, Sacramento region during the Bay Area/Sacramento period (Periods 5-6). 268 As shown in Figures 2 and 4, particles from biomass burning were elevated in the submicron 269 mode during these two period representing up to  $\sim 61\%$  of submicron particles. Biomass burning 270 particles are characterized by an intense potassium peak  $({}^{39}K^+)$  in addition to carbonaceous peaks 271 (both elemental and organic), organic nitrogen peaks (<sup>26</sup>CN<sup>-</sup>, <sup>42</sup>CNO<sup>-</sup>), ion peaks associated with 272 potassium salts (<sup>113</sup>K<sub>2</sub>Cl<sup>+</sup>, <sup>213</sup>K<sub>3</sub>SO<sub>4</sub><sup>+</sup>, etc.), and secondary species such as sulfate and/or nitrate 273 (Figure 5b) [Pratt et al., 2010; Qin and Prather, 2006; Silva et al., 1999]. 274

The key difference in particle composition observed in Northern California during the 275 Bay Area/Sacramento period (Period 6) was the high percentage of submicron OC particles 276 detected, constituting  $\sim 29 \pm 23\%$  of submicron particles by number on average and up to  $\sim 78\%$ . 277 Unlike the OC detected during the Ports of LA/LB period, which originated from fossil fuel 278 sources in the Ports of LA and LB, OC detected in Sacramento is likely the result of biogenic 279 emissions due to the fact that biogenics were dominant during the first half of the Carbonaceous 280 Aerosol Radiative Effects Study (CARES) in the Sacramento Valley [Cahill et al., 2012]. 281 Numerous new particle formation (NPF) events were observed across the Sacramento Valley 282 during CalNex and the first part of CARES during this period [Ahlm et al., 2012; Setvan et al., 283 2012]. While the lower size limit of the nozzle-inlet ATOFMS (~0.2 um) cannot be used to 284 probe the composition of newly formed particles, ATOFMS measurements can provide insight 285

286 into the composition of these particles as they grow to sizes detectable by the instrument [Creamean et al., 2011]. Figure 6 shows particle number concentration as a function of diameter 287 measured by a differential mobility particle sizer (DMPS) [Bates et al., 2012]. Two distinct 288 289 events are highlighted in black boxes when high number concentrations of small particles (~0.02 um) were observed followed by rapid growth, which is indicative of NPF events [Creamean et 290 al., 2011; Hegg and Baker, 2009; Kulmala, 2003]. As shown in Figure 6, the percentage of OC 291 particles was smallest when the NPF event was actually occurring. This is because the size of 292 the newly formed particles is below the size detection limit of the instrument. The percentage of 293 OC particles increased primarily after these NPF events occurred as the newly formed particles 294 grew to sizes detectable by the instrument suggesting that organics contributed to this growth in 295 agreement with previous studies [Creamean et al., 2011; Kulmala, 2003; Smith et al., 2008; 296 297 Zhang et al., 2004], and with recent measurements during CARES in the Sacramento Valley from an aerosol mass spectrometer (AMS) [Setvan et al., 2012]. The formation of new particles 298 occurred under conditions of lower RH and intense solar radiation; these low RH conditions also 299 300 favor the formation of high molecular weight secondary organic aerosol from biogenic emissions that may have contributed to particle formation and/or growth [Nguyen et al., 2011]. Growth of 301 these particles into sizes detectable by the ATOFMS occurred as RH increased with many of the 302 detected OC particles lacking negative ion spectra ( $85 \pm 16\%$  on average) indicating that an 303 appreciable amount of particulate water was associated with these particles [Neubauer et al., 304 305 1997; Neubauer et al., 1998].

306 3.2.2 Supermicron Particle Chemistry

307 Aged sea salt dominated the supermicron particle composition during most periods making up  $50 \pm 27\%$  of total supermicron particles on average, as shown in Figures 2 and 4. 308 Aged sea salt particles result from heterogeneous reactions with gas-phase nitrogen oxides (e.g., 309  $N_2O_{5(g)}$ , HNO<sub>3(g)</sub>), which lead to the displacement of chloride and the formation of particulate 310 nitrate [Behnke et al., 1991; Chang et al., 2011; Gard et al., 1998; Vogt et al., 1996]. In addition 311 to nitrate, methanesulfonic acid (MSA) and sulfate were also found on ~14% of aged sea salt 312 particles, on average, possibly contributing to the observed particulate chloride displacement 313 [Hopkins et al., 2008; Laskin et al., 2012]. During the Marine/Coastal Transport period (Period 314 3), fresh sea salt and marine biogenic particle types dominated the supermicron particle 315 composition making up  $33 \pm 11\%$  and  $30 \pm 17\%$  of supermicron particles on average, 316 respectively. Marine biogenic particle types include Mg-type [Gaston et al., 2011] and sulfur-317 318 type particles (C.J. Gaston et al., manuscript in preparation, 2013), which represent oceanderived particle types associated with marine biological activity and/or dissolved organic 319 material. The combined high wind speeds in addition to the elevated biological activity, as 320 321 evidenced by the presence of red tide blooms along the Southern California coast (www.sccoos.org), explains the dominance of fresh sea salt and marine biogenic emissions 322 during this period. 323

Marine particle types were negligible in the inland, Sacramento region during the Bay Area/Sacramento period. Instead, dust particles were found to represent  $39 \pm 23\%$  of supermicron particles, and biological particles were found to constitute up to ~64% of supermicron particles as shown in Figures 2 and 4f. Most of the biological particles detected by ATOFMS were determined to be spores containing dipicolinic acid, a compound that is easily detected using laser/desorption ionization at 266 nm radiation and serves as a unique matrix 15

allowing for the detection of amino acids, which typically have low absorption cross sections at 330 266 nm [Silva and Prather, 2000; Srivastava et al., 2005]. Consistent with previous 331 measurements of spores using laser/desorption ionization at 266 nm, spores detected during 332 CalNex contained  ${}^{39}K^+$ ,  ${}^{59}(CH_3)_3N^+$ ,  ${}^{74}(CH_3)_4N^+$ , m/z +86 due to either  $(C_2H_5)_2NCH_2^+$  or the 333 amino acids [leucine-HCO<sub>2</sub>] and [isoleucine-HCO<sub>2</sub>], m/z +104, which is yet to be identified, 334 phosphate ( $^{63}PO_2$ ,  $^{79}PO_3$ ), organic nitrogen ( $^{26}CN$ ,  $^{42}CNO$ ), and m/z -123, which is most likely 335 due to dipicolinic acid-HCO<sub>2</sub> [Srivastava et al., 2005] (Figure 5c). Spores have rarely been 336 detected by ATOFMS in ambient environments and represent a very unique particle type most 337 likely detected due to agricultural and biogenic emissions from the Sacramento area. 338

### 339 3.3 Variations in the Mixing-State of Carbonaceous Particle Types

#### 340 3.3.1 Soot Particle Mixing-State

In addition to probing overall particle composition, observed differences in the mixing-341 state of soot-dominated particles were also investigated (Figure 7a). These particles are 342 characterized by intense elemental carbon ion peaks  $\binom{1^2}{C^+}, \binom{3^6}{3^4}, \binom{4^8}{4^8}, \binom{4^8}{4^8}, \binom{4^8}{4^8}, \binom{4^8}{4^8}$  A small 343 percentage of freshly emitted elemental carbon particles, herein referred to as "soot", contained 344 intense elemental carbon peaks indicative of long chain elemental carbon (e.g. elemental carbon 345 ion peaks extending out to higher masses in both the positive and negative ion spectra) with only 346 low intensity peaks from secondary species such as sulfate and/or nitrate [Cahill et al., 2012; 347 Moffet and Prather, 2009]. Most of the soot-containing particles, however, contained intense 348 elemental carbon peaks that did not extend out to higher masses and were internally mixed with 349 low intensity organic peaks ( ${}^{27}C_2H_3^+$ ,  ${}^{43}C_2H_3O^+$ , etc.) in addition to other secondary species (e.g., 350 nitrate and/or sulfate) in agreement with aircraft observations during CalNex [Cahill et al., 2012; 351

352 *Metcalf et al.*, 2012]. This second class of soot-containing particles is herein referred to as Aged Soot particles, consistent with *Cahill et al.* [2012], and is further sub-divided into Aged Soot (No 353 Negatives), Aged Soot (Sulfate), Aged Soot (Nitrate), Aged Soot (Neg OC) particle classes. 354 Overall, the majority of Aged Soot particles ( $\sim 62 \pm 20\%$  of soot-containing particles on average) 355 lacked negative ion spectra and are classified as Aged Soot (No Negatives) (see Figures 5d and 356 7a), with sizes peaking in number concentration at  $\sim 0.5-0.6$  um and extending up into the 357 supermicron size range (Figure 4). The lack of negative ion spectra and larger sizes suggests that 358 these particles contained appreciable particulate water [Moffet et al., 2008b; Neubauer et al., 359 1997; Neubauer et al., 1998] possibly due to cloud or fog processing. 360

The most striking trend shown in Figure 7a is the high percentage of Aged Soot (sulfate) 361 particles detected in Southern California (representing, on average,  $\sim 30 \pm 18\%$  of detected 362 submicron soot-containing particles during the Riverside Transport, Stagnant/Ports Transport, 363 364 Marine/Coastal Transport, Ports of LA/LB period) that decreased significantly to  $\sim 11 \pm 10\%$  as the ship moved north along the California coast. Aged Soot (sulfate) particles are characterized 365 by intense sulfate peaks (<sup>80</sup>SO<sub>3</sub><sup>-</sup>, <sup>97</sup>HSO<sub>4</sub><sup>-</sup>) (Figure 5e) [*Moffet and Prather*, 2009]. This trend in 366 367 Aged Soot (sulfate) particles is in contrast to aircraft measurements made by ATOFMS, which mainly observed internal mixtures of soot and nitrate as opposed to sulfate in Southern California 368 [Cahill et al., 2012] due to the accumulation of ammonium nitrate coatings [Metcalf et al., 2012]. 369 The most likely explanation is that soot-containing particles were measured in the western port 370 regions of the LA Basin onboard the R/V Atlantis where fresh shipping and port emissions 371 prevailed as opposed to the more aged emissions sampled onboard the aircraft. Fresh emissions 372 from industrial sources at the port and ocean-going vessels contain high levels of SO<sub>2(g)</sub> leading 373 to elevated levels of particulate sulfate [Agrawal et al., 2008; Ault et al., 2010; Corbett and 374 17

Fischbeck, 1997; Corbett and Koehler, 2003]. During CalNex, Aged Soot (sulfate) particles 375 376 were typically found to peak at  $\sim 0.55$  um on average, which is larger than expected for freshly emitted soot particles [Moffet and Prather, 2009]; this is likely due to condensation and aqueous 377 phase processing of  $SO_{2(g)}$ , which has been shown to contribute high quantities of sulfate to 378 particles in this size range in the LA Basin [Hegg, 1985; Hering and Friedlander, 1982; Meng 379 and Seinfeld, 1994]. It should be noted that regulations were adopted in 2009 requiring ships to 380 switch from high sulfur to low sulfur fuel within 24 nautical miles of the California coast [CARB, 381 2009]. Hence, in addition to fresh port and shipping emissions, another potential source of 382 elevated sulfate on soot particles, particularly during Marine/Coastal Transport conditions 383 (Period 3), is the presence of red tide blooms in Southern Californian waters during CalNex 384 (www.sccoos.org) of L. polyedrum, a marine organism that has been shown to contribute 385 biogenic sulfur to aerosols [Gaston et al., 2010]. Although ATOFMS cannot be used to 386 distinguish biogenic and anthropogenic sulfate, MSA, an ocean-derived biogenic form of sulfur, 387 can be detected and used to assess the possibility of biogenic contributions to detected sulfate 388 389 levels [Gaston et al., 2010]. In fact, ~38% of Aged Soot (sulfate) particles, on average during the Riverside, Stagnant/Ports, Marine, and Ports of LA/LB periods, were found to also contain 390 MSA with the highest percentage (~62% of Aged Soot (sulfate) particles) occurring during the 391 Marine/Coastal Transport period while the lowest percentage (~13% of Aged Soot (sulfate) 392 particles) was observed during the Ports of LA/LB period. This suggests that at least some of the 393 394 observed sulfate was from a biogenic source.

Agricultural emissions were also found to impact soot mixing-state by contributing Aged Soot (nitrate) particles, particularly during the Riverside Transport and Bay Area/Sacramento periods and toward the end of the Stagnant/Ports Transport period, when air masses were 18

transported from the Central Valley. These particles had similar positive ion markers to the 398 Aged Soot (sulfate) particles, but contained nitrate peaks (<sup>46</sup>NO<sub>2</sub>, <sup>62</sup>NO<sub>3</sub>) that were more intense 399 than sulfate (see Figure 5f). This particle type has been shown to result from the acquisition of 400 nitrate namely formed through photo-chemically produced nitric acid (HNO<sub>3(0)</sub>), which 401 condenses onto particles causing them to grow; aqueous phase processing may also contribute to 402 this growth and the acquisition of nitrate [Moffet and Prather, 2009; Moffet et al., 2008b]. This 403 particle type represented up to  $\sim 33\%$  of soot-containing particles and was observed to peak at a 404 larger size mode than the other soot-containing particle types at  $\sim 0.75$  um during the campaign 405 406 (Figure 4).

In addition to Aged Soot (nitrate) particles, unique Aged Soot particles containing sulfate 407 and intense ions at m/z -43, -57, and -71 (see Figure 5g) were detected during a period of 408 agricultural influence (Inland/Valley Transport conditions). To the best of our knowledge, 409 410 mixtures of soot and these ions have never been detected before by ATOFMS. Since organic peaks in Aged Soot particles typically appear as positive ions, this particle type has been labeled 411 Aged Soot (Neg OC). It is most likely that this particle type is derived from a unique, fresh 412 413 source near the sampling site, since these ion markers were not observed on other particle types. Further evidence for this speculation comes from the fact that this particle type peaks at a smaller 414 size (~0.35 um) than the other soot-containing types, as shown in Figure 4e, suggesting a local 415 source. The unique organic markers likely correspond to  ${}^{43}CH_3CO^2$ ,  ${}^{57}C_2O_2H^2$ , and  ${}^{71}C_3H_3O_2^2$ 416 [McLafferty and Turecek, 1993; Silva et al., 1999; Silva and Prather, 2000] possibly due to 417 contributions from levoglucason and/or methylglyoxal (m/z -71), glyoxal (m/z -57), and 418 acetaldehyde (m/z -43) [Silva et al., 1999; Silva and Prather, 2000]. These organic species could 419 represent secondary organic aerosol (SOA) formed from the photolytic processing of organics in 420 19

the aqueous phase [*Bateman et al.*, 2011]; this mechanism was suggested to contribute organic
components with elevated atomic O:C ratios during CalNex [*Duong et al.*, 2011]. Additional
field and laboratory measurements are required to confirm the identification of these ion peaks.

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# 3.3.2 The Mixing-State of Organic Particles

In addition to soot-containing particles, the mixing-state of OC particles (e.g., non-soot 425 426 containing OC particles) was also investigated by examining temporal trends. These non-soot containing OC particles are distinguished from the Aged Soot types described above in that these 427 particles are not characterized by intense elemental carbon peaks at m/z + 12, +24, +36, etc., but 428 instead have m/z + 27, +37, and/or +43 as the main positive ion peaks. Three types of OC 429 particles were identified during CalNex: OC (no negatives), OC (sulfate), and OC (nitrate). 430 Temporal trends of the OC particle types are shown in Figure 7b. Most OC particles were found 431 to lack negative ion spectra ( $\sim 67 \pm 37\%$  on average) (see Figure 5h) and peaked in the 0.5-0.6 432 um size range (Figure 4) suggesting that they contained appreciable particulate water similar to 433 434 soot-containing particles. However, during Marine/Coastal and Inland/Valley Transport conditions (Periods 3 and 5), OC particles containing intense sulfate peaks (see Figure 5i) were 435 more common, representing  $\sim 55 \pm 40\%$  of organic particles on average, likely due to photo-436 chemically produced sulfate possibly derived from biogenic emissions. In addition to sulfate, 437 OC particles with intense nitrate peaks (see Figure 5j) were also observed namely at night during 438 the Marine/Coastal Transport, Ports of LA/LB, Inland/Valley Transport, and Bay 439 Area/Sacramento periods (Periods 3-6) and represented  $\sim 9 \pm 17\%$  of organic particles, on 440 average. Overall, OC (nitrate) particles peaked at a smaller particle size (~0.35 um) suggesting 441 442 that not all of the nitrate on these particles was due to photo-chemically-produced nitrate, which typically leads to larger particle sizes. Instead, these observations suggest possible contributions of organonitrates formed from reactions with gas-phase precursors and nitrate radical at night that then condense onto particles [*Ng et al.*, 2008]; organonitrates were also found to contribute to the organic aerosol burden during CalNex in Bakersfield, CA [*Rollins et al.*, 2012].

The OC types described above also frequently contained aromatic peaks  $({}^{51}C_4H_3^+,$ 447 <sup>63</sup>C<sub>5</sub>H<sub>3</sub><sup>+</sup>, <sup>77</sup>C<sub>6</sub>H<sub>5</sub><sup>+</sup>, etc) [Silva and Prather, 2000], which have been associated with secondary 448 processing of organic species in vehicle exhaust [Shields et al., 2007; Sodeman et al., 2005; 449 Spencer et al., 2006; Toner et al., 2008] and humic substances formed from biomass burning 450 [Holecek et al., 2007; Mavol-Bracero et al., 2002; Qin and Prather, 2006]. OC particles also 451 contained ion peaks indicative of amines (e.g.,  ${}^{59}(CH_3)_3N^+$ ,  ${}^{86}(C_2H_5)_2NCH_2^+$ ,  ${}^{101}(C_2H_5)_3N^+$ , 452 <sup>118</sup>(C<sub>2</sub>H<sub>5</sub>)<sub>3</sub>NOH<sup>+</sup>, etc.), which are semi-volatile species that can partition onto pre-existing 453 particles [Angelino et al., 2001; Pratt et al., 2009; Schade and Crutzen, 1995], and oxidized 454 organic markers (e.g.,  ${}^{43}C_2H_3O^+$ ) that indicate the presence of SOA [*Oin et al.*, 2012] (see Figure 455 5h). Ternary plots were used to examine the prevalence of these compounds and to further 456 elucidate the mixing-state and sources of OC particles using the ion peaks  ${}^{59}(CH_3)_3N^+$ ,  ${}^{43}C_2H_3O^+$ , 457 and  ${}^{77}C_6H_5^+$  as markers for amines, SOA/oxidized organics, and aromatics, respectively (see 458 Figure 8). 459

Amines were found to dominate the OC mixing-state during the first two periods, with 65  $\pm 33\%$  of OC particles containing intense amine markers during the Riverside Transport period and  $48 \pm 44\%$  during the Stagnant/Ports Transport period (Figure 8). Dairy farm emissions from the Chino and Central Valley regions in addition to possible contributions from traffic emissions wherein ammonia and amines can be produced from catalytic converters [*Fraser and Cass*, 465 1998; Sodeman et al., 2005], explain the dominance of amines during these two time periods [Hughes et al., 2002; Pastor et al., 2003; Pratt et al., 2009; Oin et al., 2012; Schade and 466 Crutzen, 1995; Sorooshian et al., 2008]. Further, during these two periods, low temperatures 467 468 and the highest average RH values were measured on the ship; similar meteorological conditions encountered during transport would also favor the detection of amines during these periods 469 [Angelino et al., 2001]. Further,  $18 \pm 23\%$  of OC particles contained intense amine markers 470 during the Ports of LA/LB period even though inland transport conditions were not encountered 471 suggesting that ports could contribute an industrial source of amines, possibly from vehicle 472 emissions including heavy duty diesel vehicles [Bishop et al., 2012; Fraser and Cass, 1998; 473 Sodeman et al., 2005]. Amines could also have resulted from marine biogenic emissions 474 [Facchini et al., 2008a; Sorooshian et al., 2009]. 475

Oxidized organics were prevalent during Inland/Valley Transport conditions (Period 5) 476 477 (Figure 8e). The high  $O_3$  concentrations and photo-chemically aged nature of the sampled air masses suggest that the organics during the Inland/Valley Transport period resulted from 478 secondary rather than primary sources [Na et al., 2004; Oin et al., 2012]. Oxidized organics at 479 480 m/z +43 were also prevalent during the Stagnant/Ports Transport period, when photo-chemically aged air masses were sampled, and dominant during Marine/Coastal Transport conditions 481 (Period 3), likely due to secondary contributions. Another source of the oxidized organics during 482 the Marine/Coastal Transport period could be organics from marine biogenic sources, such as 483 lipopolysaccharides, that have a higher oxygen content than organics from anthropogenic sources 484 485 [Facchini et al., 2008b; Ovadnevaite et al., 2011; Russell et al., 2011; Russell et al., 2010].

486 SOA from aromatics was more prevalent during the Ports of LA/LB and Bay Area/Sacramento periods than during any other periods; however, oxidized organics were still 487 the most prevalent type of organic during these two periods. The increased frequency of 488 489 aromatic markers in the Ports of LA and LB is likely due to increased emissions from diesel combustion by trucks [Kasper et al., 2007; Maricq, 2007; Shields et al., 2007; Spencer et al., 490 2006] in addition to emissions from ships [Kasper et al., 2007; Lack et al., 2009; Murphy et al., 491 2009; Russell et al., 2009]. Oxidized organics and amines in Sacramento are likely from 492 agricultural and biogenic emissions [Chow et al., 2006b; Setyan et al., 2012; Sorooshian et al., 493 2008] while the observed SOA from aromatics are likely humic substances derived from biomass 494 burning [Holecek et al., 2007; Mayol-Bracero et al., 2002]. 495

# 496 4. Conclusions and Implications for California

Overall, we found that the chemical properties of aerosol particles differed widely across 497 California based on particle source and transport conditions. As evidenced by the particle 498 499 chemistry and mixing-state measured onboard the R/V Atlantis, Southern California was impacted by fresh shipping and vehicular emissions from cars and diesel trucks in addition to 500 marine biogenic emissions. Soot-containing particles were the most prevalent submicron 501 particles in Southern California and were typically mixed with sulfate while aromatic markers 502 were found to contribute to organics due to residual fuel, distillate fuel, and gasoline combustion. 503 These fresh emissions were episodically mixed with inland, agricultural emissions leading to the 504 presence of secondary species such as amines. It is important to note that the particles measured 505 in Southern California contained elevated levels of sulfate because these measurements focused 506 507 on fresh emissions from the Ports of LA and LB; if these measurements were made further

inland of the ports, enhancements in nitrate namely due to particle aging and contributions of
ammonium nitrate, as seen by other measurements, would have been observed instead [*Cahill et al.*, 2012; *Metcalf et al.*, 2012].

In contrast, Northern California was impacted by biogenic emissions from the forested Sierra Nevada foothills and agricultural emissions as evidenced by the prevalence of organic carbon [*Cahill et al.*, 2012; *Setyan et al.*, 2012]. Oxidized organic compounds were frequently associated with particles during Inland/Valley transport conditions and in Sacramento, likely due to contributions from SOA while aromatic markers observed in Sacramento are likely due to contributions from biomass burning.

517 Overall, we found that the mixing-state of particles in California varies significantly on a 518 regional scale due to diverse local sources. Further, the mixing-state of particles from these local 519 emissions can be significantly altered based on differences in meteorological conditions and air 520 mass histories, which will most likely result in regional differences in the health, optical, and 521 cloud nucleating properties for the aerosol populations observed across California. This should 522 be taken into account when determining which emissions sources to regulate in order to mitigate 523 the adverse effects of aerosols on both human health and climate change.

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## 540 **References:**

- Agrawal, H., Q.G.J. Malloy, W.A. Welch, J.W. Miller, and D.R. Cocker (2008), In-use gaseous
   and particulate matter emissions from a modern ocean going container vessel, *Atmos. Environ.*, 42 (21), 5504-5510.
- Ahlm, L., S. Liu, D.A. Day, L.M. Russell, R. Weber, D.R. Gentner, A.H. Goldstein, J.P. 544 DiGangi, S.B. Henry, F.N. Keutsch, T.C. VandenBoer, M.Z. Markovic, J.G. Murphy, X. 545 Ren, and S. Scheller (2012), Formation and growth of ultrafine particles from secondary 546 547 sources in Bakersfield. California, JGeophys Res. 117. D00V08. doi:10.1029/2011JD017144. 548
- Allen, J.O. (2002), YAADA software toolkit to analyze single-particle mass spectral data:
   Reference manual version 1.1, *Arizona State University*, <u>http://www.yaada.org</u>.
- Angelino, S., D.T. Suess, and K.A. Prather (2001), Formation of aerosol particles from reactions
   of secondary and tertiary alkylamines: Characterization by aerosol time-of-flight mass
   spectrometry, *Environ. Sci. Tech.*, *35* (15), 3130-3138.
- Ault, A.P., C.J. Gaston, Y. Wang, G. Dominguez, M.H. Thiemens, and K.A. Prather (2010),
   Characterization of the single particle mixing state of individual ship plume events
   measured at the Port of Los Angeles, *Environ. Sci. Tech.*, 44 (6), 1954-1961.
- Ault, A.P., M.J. Moore, H. Furutani, and K.A. Prather (2009), Impact of emissions from the Los
   Angeles port region on San Diego air quality during regional transport events, *Environ. Sci. Tech.*, 43 (10), 3500-3506.
- Bateman, A.P., S.A. Nizkorodov, J. Laskin, and A. Laskin (2011), Photolytic processing of
   secondary organic aerosols dissolved in cloud droplets, *Phys. Chem. Chem. Phys.*, 13,
   12199-12212.
- Bates, T.S., P.K. Quinn, D.J. Coffman, D.S. Covert, T.L. Miller, J.E. Johnson, G.R. Carmichael,
  I. Uno, S.A. Guazzotti, D.A. Sodeman, K.A. Prather, M. Rivera, L.M. Russell, and J.T.
  Merrill (2004), Marine boundary layer dust and pollutant transport associated with the
  passage of a frontal system over eastern Asia, *J. Geophys. Res.-[Atmos.]*, 109 (D19S19),
  doi:10.1029/2003JD004094.
- Bates, T.S., P.K. Quinn, A.A. Frossard, L.M. Russell, J. Hakala, T. Petaja, M. Kulmala, D.S.
  Covert, C.D. Cappa, S.-M. Li, K. Hayden, I. Nuaaman, R. McLaren, P. Massoli, M.R.
  Canagaratna, T.B. Onasch, D. Sueper, D. Worsnop, and W.C. Keene (2012),
  Measurements of ocean derived aerosol off the coast of California, *J Geophys Res*, *117*,
  D00V15, doi:10.1029/2012JD017588.
- Behnke, W., H.U. Kruger, V. Scheer, and C. Zetzsch (1991), Formation of Atomic Cl from Sea
   Spray Via Photolysis of Nitryl Chloride Determination of the Sticking Coefficient of
   N<sub>2</sub>O<sub>5</sub> on NaCl Aerosol, *J. Aerosol Sci.*, 22, S609-S612.
- Bishop, G.A., B.G. Schuchmann, and D.H. Stedman (2012), Emission changes resulting from the
  San Pedro Bay, California ports truck retirement program, *Environ. Sci. Tech.*, 46, 551558.
- Blanchard, D.C., and A.H. Woodcock (1957), Bubble formation and modification in the sea and
  its meteorological significance, *Tellus*, 9 (2), 145-158.
- Bon, D.M., I.M. Ulbrich, J.A. de Gouw, C. Warneke, W.C. Kuster, M.L. Alexander, A. Baker,
  A.J. Beyersdorf, D. Blake, R. Fall, J.L. Jimenez, S.C. Herndon, L.G. Huey, W.B.
  Knighton, J. Ortega, S. Springston, and O. Vargas (2011), Measurements of volatile

- organic compounds at a suburban ground site (T1) in Mexico City during the MILAGRO
  2006 campaign: measurement comparison, emission ratios, and source attribution, *Atmos. Chem. Phys.*, 11, 2399-2421.
- Cahill, J.F., K. Suski, J. Seinfeld, R.A. Zaveri, and K.A. Prather (2012), The mixing state
   between northern and southern California measured during CARES and CalNEX 2010,
   *Atmos. Chem. Phys.*, *12*, 10989-11002, doi:10.5194/acp-12-10989-2012.
- CARB, Final Regulation Order. Fuel Sulfur and Other Operational Requirements for Ocean Going Vessels within California Waters and 24 Nautical Miles of the California Baseline,
   California Air Resources Board: Sacramento, CA, 2009.
- Chang, W.L., P.V. Bhave, S.S. Brown, N. Riemer, J. Stutz, and D. Dabdub (2011),
   Heterogeneous atmospheric chemistry, ambient measurements, and model calculations of
   N<sub>2</sub>O<sub>5</sub>: A review, *Aerosol Sci. Tech.*, 45 (6), 665-695.
- Chen, L.W.A., J.G. Watson, J.C. Chow, and K.L. Magliano (2007), Quantifying PM<sub>2.5</sub> source
   contributions for the San Joaquin Valley with multivariate receptor models, *Environ. Sci. Tech.*, 41 (8), 2818-2826.
- Chow, J.C., L.W.A. Chen, J.G. Watson, D.H. Lowenthal, K.A. Magliano, K. Turkiewicz, and
   D.E. Lehrman (2006a), PM<sub>2.5</sub> chemical composition and spatiotemporal variability during
   the California Regional PM<sub>10</sub>/PM<sub>2.5</sub> Air Quality Study (CRPAQS), *J. Geophys. Res.- [Atmos.]*, 111 (D10), D10S04, doi:10.1029/2005JD006457.
- Chow, J.C., J.G. Watson, D.H. Lowenthal, L.W.A. Chen, and K.L. Magliano (2006b),
   Particulate carbon measurements in California's San Joaquin Valley, *Chemosphere*, 62 (3), 337-348.
- 606 Corbett, J.J., and P. Fischbeck (1997), Emissions from ships, *Science*, 278 (5339), 823-824.
- 607 Corbett, J.J., and H.W. Koehler (2003), Updated emissions from ocean shipping, *J. Geophys.* 608 *Res.-[Atmos.]*, 108 (D20), 4650, doi:10.1029/2003JD003751.
- Creamean, J.M., A.P. Ault, J.E. Ten Hoeve, M.Z. Jacobson, G.C. Roberts, and K.A. Prather
  (2011), Measurements of aerosol chemistry during new particle formation events at a
  remote rural mountain site, *Environ. Sci. Tech.*, 45 (19), 8208-8216.
- Dall'Osto, M., R.M. Harrison, D.C.S. Beddows, E.J. Freney, M.R. Heal, and R.J. Donovan
   (2006), Single-particle detection efficiencies of aerosol time-of-flight mass spectrometry
   during the North Atlantic marine boundary layer experiment, *Environ. Sci. Tech.*, 40 (16),
   5029-5035.
- Docherty, K.S., E.A. Stone, I.M. Ulbrich, P.F. DeCarlo, D.C. Snyder, J.J. Schauer, R.E. Peltier,
  R.J. Weber, S.M. Murphy, J.H. Seinfeld, B.D. Grover, D.J. Eatough, and J.L. Jimenez
  (2008), Apportionment of primary and secondary organic aerosols in Southern California
  during the 2005 Study of Organic Aerosols in Riverside (SOAR-1), *Environ. Sci. Tech.*,
  42 (20), 7655-7662.
- Draxler, R.R., and G.D. Rolph (2011), HYSPLIT (HYbrid Single-Particle Lagrangian Integrated
   Trajectory) Model access via NOAA ARL READY Website
   (<u>http://ready.arl.noaa.gov/HYSPLIT.php</u>), NOAA Air Resources Laboratory, Silver
   Spring, MD.
- Duong, H.T., A. Sorooshian, J.S. Craven, S.P. Hersey, A.R. Metcalf, X. Zhang, R.J. Weber, H.H.
   Jonsson, R.C. Flagan, and J.H. Seinfeld (2011), Water-soluble organic aerosol in the Los
   Angeles Basin and outflow regions: Airborne and ground measurements during the 2010
   CalNex field campaign, *J. Geophys. Res.*, *116*, D00V04, doi:10.1029/2011JD016674.

- Eatough, D.J., B.D. Grover, W.R. Woolwine, N.L. Eatough, R. Long, and R. Farber (2008),
  Source apportionment of 1 h semi-continuous data during the 2005 Study of Organic
  Aerosols in Riverside (SOAR) using positive matrix factorization, *Atmos. Environ.*, 42
  (11), 2706-2719.
- Facchini, M.C., S. Decesari, M. Rinaldi, C. Carbone, E. Finessi, M. Mircea, S. Fuzzi, F. Moretti,
  E. Tagliavini, D. Ceburnis, and C.D. O'Dowd (2008a), Important source of marine
  secondary organic aerosol from biogenic amines, *Environ. Sci. Tech.*, 42 (24), 91169121.
- Facchini, M.C., M. Rinaldi, S. Decesari, C. Carbone, E. Finessi, M. Mircea, S. Fuzzi, D.
  Ceburnis, R. Flanagan, E.D. Nilsson, G. de Leeuw, M. Martino, J. Woeltjen, and C.D.
  O'Dowd (2008b), Primary submicron marine aerosol dominated by insoluble organic
  colloids and aggregates, *Geophys. Res. Lett.*, 35, L17814, doi:10.1029/2008GL034210.
- Fitzgerald, J.W. (1991), Marine aerosols: A review, *Atmos. Environ. A-Gen. Topics*, 25 (3-4),
   533-545.
- Fraser, M.P., and G.R. Cass (1998), Detection of excess ammonia emissions from in-use vehicles
  and the implications for fine particle control, *Environ. Sci. Tech.*, *32*, 1053-1057.
- Gard, E., J.E. Mayer, B.D. Morrical, T. Dienes, D.P. Fergenson, and K.A. Prather (1997), Realtime analysis of individual atmospheric aerosol particles: Design and performance of a
  portable ATOFMS, *Anal. Chem.*, *69* (20), 4083-4091.
- Gard, E.E., M.J. Kleeman, D.S. Gross, L.S. Hughes, J.O. Allen, B.D. Morrical, D.P. Fergenson,
  T. Dienes, M.E. Galli, R.J. Johnson, G.R. Cass, and K.A. Prather (1998), Direct
  observation of heterogeneous chemistry in the atmosphere, *Science*, 279 (5354), 11841187.
- Gaston, C.J., H. Furutani, S.A. Guazzotti, K.R. Coffee, T.S. Bates, P.K. Quinn, L.I. Aluwihare,
  B.G. Mitchell, and K.A. Prather (2011), Unique ocean-derived particles serve as a proxy
  for changes in ocean chemistry, *J. Geophys. Res.-[Atmos.]*, *116*, D18310,
  doi:10.1029/2010JD015289.
- Gaston, C.J., K.A. Pratt, X.Y. Qin, and K.A. Prather (2010), Real-time detection and mixing
  state of methanesulfonate in single particles at an inland urban location during a
  phytoplankton bloom, *Environ. Sci. Tech.*, 44 (5), 1566-1572.
- Gelencser, A., K. Siszler, and J. Hlavay (1997), Toluene-benzene concentration ratio as a tool for
   characterizing the distance from vehicular emission sources, *Environ. Sci. Tech.*, *31*,
   2869-2872.
- Grover, B.D., N.L. Eatough, W.R. Woolwine, J.P. Cannon, D.J. Eatough, and R.W. Long (2008),
   Semi-continuous mass closure of the major components of fine particulate matter in
   Riverside, CA, *Atmos. Environ.*, 42, 250-260.
- Guazzotti, S.A., K.R. Coffee, and K.A. Prather (2001), Continuous measurements of size resolved particle chemistry during INDOEX-Intensive Field Phase 99, *J. Geophys. Res. [Atmos.]*, 106 (D22), 28607-28627.
- Hawkins, L.N., L.M. Russell, D.S. Covert, P.K. Quinn, and T.S. Bates (2010), Carboxylic acids,
  sulfates, and organosulfates in processed continental organic aerosol over the southeast
  Pacific Ocean during VOCALS-REx 2008, *J. Geophys. Res.-[Atmos.]*, *115*, D13201,
  doi:10.1029/2009JD013276.

- Healy, R.M., I.P. O'Connor, S. Hellebust, A. Allanic, J.R. Sodeau, and J.C. Wenger (2009),
  Characterization of single particles from in-port ship emissions, *Atmos Environ*, 43,
  674 6408-6414.
- Hegg, D.A. (1985), The importance of liquid-phase oxidation of SO<sub>2</sub> in the troposphere, J. *Geophys. Res.*, 90 (D2), 3773-3779.
- Hegg, D.A., and M.B. Baker (2009), Nucleation in the atmosphere, *Reports on Progress in Physics*, 72 (5), doi:10.1088/0034-4885/72/5/056801.
- Hering, S.V., and S.K. Friedlander (1982), Origins of aerosol sulfur size distributions in the Los
   Angeles Basin, *Atmos Environ*, 16 (11), 2647-2656.
- 681 Holecek, J.C., M.T. Spencer, and K.A. Prather (2007), Analysis of rainwater samples: Comparison of single particle residues with ambient particle chemistry from the northeast 682 Pacific Indian oceans, Res.-[Atmos.], 683 and J. Geophys. 112 (D22), doi:10.1029/2006JD008269. 684
- Hopkins, R.J., Y. Desyaterik, A.V. Tivanski, R.A. Zaveri, C.M. Berkowitz, T. Tyliszczak, M.K. 685 Gilles, and A. Laskin (2008), Chemical speciation of sulfur in marine cloud droplets and 686 particles: Analysis of individual particles from the marine boundary layer over the 687 688 California Current, J. Geophys. Res.-[Atmos.], 113 (D4), D04209, doi:10.1029/2007JD008954. 689
- Hughes, L.S., J.O. Allen, P. Bhave, M.J. Kleeman, G.R. Cass, D.Y. Liu, D.F. Fergenson, B.D.
   Morrical, and K.A. Prather (2000), Evolution of atmospheric particles along trajectories
   crossing the Los Angeles basin, *Environ. Sci. Tech.*, *34* (15), 3058-3068.
- Hughes, L.S., J.O. Allen, L.G. Salmon, P.R. Mayo, R.J. Johnson, and G.R. Cass (2002),
   Evolution of nitrogen species air pollutants along trajectories crossing the Los Angeles
   area, *Environ. Sci. Tech.*, 36 (18), 3928-3935.
- Kasper, A., S. Aufdenblatten, A. Forss, M. Mohr, and H. Burtscher (2007), Particulate emissions
  from a low-speed marine diesel engine, *Aerosol Sci. Tech.*, 41 (1), 24-32.
- 698 Kulmala, M. (2003), How particles nucleate and grow, *Science*, *302* (5647), 1000-1001.
- Lack, D.A., J.J. Corbett, T. Onasch, B. Lerner, P. Massoli, P.K. Quinn, T.S. Bates, D.S. Covert,
  D. Coffman, B. Sierau, S. Herndon, J. Allan, T. Baynard, E. Lovejoy, A.R. Ravishankara,
  and E. Williams (2009), Particulate emissions from commercial shipping: Chemical,
  physical, and optical properties, *J. Geophys. Res.-[Atmos.]*, *114*, D00F04,
  doi:10.1029/2008JD011300.
- Laskin, A., R.C. Moffet, M.K. Gilles, J.D. Fast, R.A. Zaveri, B. Wang, P. Nigge, and J.
   Shutthanandan (2012), Tropospheric chemistry of internally mixed sea salt and organic
   particles: Surprising reactivity of NaCl with weak organic acids, *J. Geophys. Res.*, 117,
   D15302, doi:10.1029/2012JD017743.
- Liu, D.-Y., K.A. Prather, and S.V. Herring (2000), Variations in the size and chemical composition of nitrate-containing particles in Riverside, CA, *Aerosol Sci. Tech.*, 33 (1-2), 710
   71-86.
- Magliano, K.L., V.M. Hughes, L.R. Chinkin, D.L. Coe, T.L. Haste, N. Kumar, and F.W.
   Lurmann (1999), Spatial and temporal variations in PM<sub>10</sub> and PM<sub>2.5</sub> source contributions
   and comparison to emissions during the 1995 integrated monitoring study, *Atmos. Environ.*, 33 (29), 4757-4773.
- Maricq, M.M. (2007), Chemical characterization of particulate emissions from diesel engines: A review, *J. Aerosol Sci.*, *38* (11), 1079-1118.

- Mayol-Bracero, O.L., P. Guyon, B. Graham, G. Roberts, M.O. Andreae, S. Decesari, M.C.
  Facchini, S. Fuzzi, and P. Artaxo (2002), Water-soluble organic compounds in biomass
  burning aerosols over Amazonia 2. Apportionment of the chemical composition and
  importance of the polyacidic fraction, *J. Geophys. Res.-[Atmos.]*, *107* (D20), 8091,
  doi:10.1029/2001JD000522.
- McLafferty, F.W., and F. Turecek, *Interpretation of Mass Spetra*, University Science Books,
   Sausalito, CA, 1993.
- Meng, Z., and J.H. Seinfeld (1994), On the source of the submicrometer droplet mode of urban
   and regional aerosols, *Aerosol Sci. Tech.*, 20, 253-265.
- 726 Metcalf, A.R., J.S. Craven, J.J. Ensberg, J. Brioude, W. Angevine, A. Sorooshian, H.T. Duong, H.H. Jonsson, R.C. Flagan, and J.H. Seinfeld (2012), Black carbon aerosol over the Los 727 during CalNex, Geophys. 728 Angeles Basin  $J_{\cdot}$ Res., 117, D00V13, doi:10.1029/2011JD017255. 729
- Moffet, R.C., Y. Desyaterik, R.J. Hopkins, A.V. Tivanski, M.K. Gilles, Y. Wang, V.
  Shutthanandan, L.T. Molina, R.G. Abraham, K.S. Johnson, V. Mugica, M.J. Molina, A.
  Laskin, and K.A. Prather (2008a), Characterization of aerosols containing Zn, Pb, and Cl
  from an industrial region of Mexico City, *Environ. Sci. Tech.*, *42* (19), 7091-7097.
- Moffet, R.C., and K.A. Prather (2009), In-situ measurements of the mixing state and optical
   properties of soot with implications for radiative forcing estimates, *PNAS*, *106* (29),
   11872-11877.
- Moffet, R.C., X. Qin, T. Rebotier, H. Furutani, and K.A. Prather (2008b), Chemically segregated
   optical and microphysical properties of ambient aerosols measured in a single-particle
   mass spectrometer, J. Geophys. Res., 113 (D12213), D12213,
   doi:10.1029/2007JD009393.
- Monahan, E.C., C.W. Fairall, K.L. Davidson, and P.J. Boyle (1983), Observed interrelations
   between 10m winds, ocean whitecaps and marine aerosols, *Quarterly Journal of the Royal Meteorological Society*, *109* (460), 379-392.
- Murphy, S.M.A., H.; Sorooshian, A.; Padro, L. T.; Gates, S.W. H.; Hersey, W. A.; Jung, H.;
  Miller, J. W.; Cocker, D. R., and A.J. Nenes, H. H.; Flagan, R. C.; Seinfeld, J. H. (2009),
  Comprehensive simultaneous shipboard and airborne characterization of exhaust from a
  modern container ship at sea., *Environ. Sci. Tech.*, *43* (13), 4626-4640.
- Na, K., A.A. Sawant, C. Song, and D.R. Cocker (2004), Primary and secondary carbonaceous
  species in the atmosphere of Western Riverside County, California, *Atmos Environ*, *38*, 1345-1355.
- Neubauer, K.R., M.V. Johnston, and A.S. Wexler (1997), On-line analysis of aqueous aerosols
   by laser desorption ionization, *International Journal of Mass Spectrometry and Ion Processes*, 163 (1-2), 29-37.
- Neubauer, K.R., M.V. Johnston, and A.S. Wexler (1998), Humidity effects on the mass spectra of single aerosol particles, *Atmos. Environ.*, *32* (14-15), 2521-2529.
- Ng, N.L., A.J. Kwan, J.D. Surratt, A.W.H. Chan, P.S. Chhabra, A. Sorooshian, H.O.T. Pye, J.D.
  Crounse, P.O. Wennberg, R.C. Flagan, and J.H. Seinfeld (2008), Secondary organic aerosol (SOA) formation from reaction of isoprene with nitrate radicals (NO<sub>3</sub>), *Atmos. Chem. Phys.*, 8 (14), 4117-4140.

- Nguyen, T.B., P.J. Roach, J. Laskin, A. Laskin, and S.A. Nizkorodov (2011), Effect of humidity
   on the composition of isoprene photooxidation secondary organic aerosol, *Atmos. Chem. Phys.*, *11*, 6931-6944.
- Noble, C.A., and K.A. Prather (1996), Real-time measurement of correlated size and
   composition profiles of individual atmospheric aerosol particles, *Environ. Sci. Tech.*, 30
   (9), 2667-2680.
- O'Dowd, C.D., and G. De Leeuw (2007), Marine aerosol production: a review of the current knowledge, *Phil. Trans. A*, 365 (1856), 1753-1774.
- Ovadnevaite, J., C. O'Dowd, M. Dall'Osto, D. Ceburnis, D.R. Worsnop, and H. Berresheim
   (2011), Detecting high contributions of primary organic matter to marine aerosol: A case
   study, *Geophys. Res. Lett.*, *38*, L02807, doi:10.1029/2010GL046083.
- Pastor, S.H., J.O. Allen, L.S. Hughes, P. Bhave, G.R. Cass, and K.A. Prather (2003), Ambient
   single particle analysis in Riverside, California by aerosol time-of-flight mass
   spectrometry during the SCOS97-NARSTO, *Atmos. Environ.*, *37*, S239-S258.
- Poschl, U. (2005), Atmospheric aerosols: Composition, transformation, climate and health
   effects, *Angewandte Chemie-International Edition*, 44 (46), 7520-7540.
- Prather, K.A., T. Nordmeyer, and K. Salt (1994), Real-Time Characterization of Individual
   Aerosol-Particles Using Time-of-Flight Mass-Spectrometry, *Anal. Chem.*, 66 (9), 1403 1407.
- Pratt, K.A., L.E. Hatch, and K.A. Prather (2009), Seasonal volatility dependence of ambient
   particle phase amines, *Environ. Sci. Tech.*, *43* (14), 5276-5281.
- Pratt, K.A., A.J. Heymsfield, C.H. Twohy, S.M. Murphy, P.J. DeMott, J.G. Hudson, R.
  Subramanian, Z.E. Wang, J.H. Seinfeld, and K.A. Prather (2010), In Situ Chemical
  Characterization of Aged Biomass-Burning Aerosols Impacting Cold Wave Clouds, J. *Atmos. Sci.*, 67 (8), 2451-2468.
- Pratt, K.A., and K.A. Prather (2009), Real-time, single-particle volatility, size, and chemical composition measurements of aged urban aerosols, *Environ. Sci. Tech.*, 43 (21), 8276-8282.
- Pratt, K.A., and K.A. Prather (2011), Mass spectrometry of atmospheric aerosols—Recent developments and applications. Part II: On-line mass spectrometry techniques, *Mass Spectrometry Reviews*, DOI 10.1002/mas.20330.
- Qin, X., K.A. Pratt, L.G. Shields, S.M. Toner, and K.A. Prather (2012), Seasonal comparisons of
   single-particle chemical mixing state in Riverside, CA, *Atmos. Environ.*, 59, 587-596.
- Qin, X.Y., P.V. Bhave, and K.A. Prather (2006), Comparison of two methods for obtaining
   quantitative mass concentrations from aerosol time-of-flight mass spectrometry
   measurements, *Anal. Chem.*, 78 (17), 6169-6178.
- Qin, X.Y., and K.A. Prather (2006), Impact of biomass emissions on particle chemistry during
   the California Regional Particulate Air Quality Study, *International Journal of Mass Spectrometry*, 258 (1-3), 142-150.
- Quinn, P.K., and T.S. Bates (2005), Regional aerosol properties: Comparisons of boundary layer
   measurements from ACE 1, ACE 2, Aerosols99, INDOEX, ACE Asia, TARFOX, and
   NEAQS, *J. Geophys. Res.*, *110* (D14202), doi:10.1029/2004JD004755.
- Roberts, J.M., F.C. Fehsenfeld, S.C. Liu, M.J. Bollinger, C. Hahn, D.L. Albritton, and R.E.
   Sievers (1984), Measurements of aromatic hydrocarbon ratios and NO<sub>x</sub> concentrations in

- the rural troposphere: Observation of air mass photochemical aging and NO<sub>x</sub> removal, Atmos Environ, 18 (11), 2421-2432.
- Rollins, A.W., E.C. Browne, K.-E. Min, S.E. Pusede, P.J. Wooldridge, D.R. Gentner, A.H.
   Goldstein, S. Liu, D.A. Day, L.M. Russell, and R.C. Cohen (2012), Evidence for NO<sub>x</sub>
   control over nighttime SOA formation, *Science*, *337* (6099), 1210-1212.
- Russell, L.M., R. Bahadur, and P.J. Ziemann (2011), Identifying organic aerosol sources by
   comparing functional group composition in chamber and atmospheric particles, *PNAS*,
   doi:10.1073/pnas.1006461108.
- Russell, L.M., L.N. Hawkins, A.A. Frossard, P.K. Quinn, and T.S. Bates (2010), Carbohydratelike composition of submicron atmospheric particles and their production from ocean
  bubble bursting, *PNAS*, 107 (15), 6652-6657.
- Russell, L.M., K.J. Noone, R.J. Ferek, R.A. Pockalny, R.C. Flagan, and J.H. Seinfeld (2000),
   Combustion organic aerosol as cloud condensation nuclei in ship tracks, *American Meteorological Society*, 2591-2606.
- Russell, L.M., S. Takahama, S. Liu, L.N. Hawkins, D.S. Covert, P.K. Quinn, and T.S. Bates 818 (2009), Oxygenated fraction and mass of organic aerosol from direct emission and 819 820 atmospheric processing measured on the R/V Ronald Brown during **TEXAQS/GoMACCS** 2006, Geophys.  $J_{\cdot}$ Res.-[Atmos.], 114, D00F05, 821 822 doi:10.1029/2008JD011275.
- Ryerson, T.B., A.E. Andrews, W.M. Angevine, T.S. Bates, C.A. Brock, B. Cairns, R.C. Cohen, 823 O.R. Cooper, J.A. de Gouw, F.C. Fehsenfeld, R.A. Ferrare, M.L. Fischer, R.C. Flagan, 824 A.H. Goldstein, J.W. Hair, R.M. Hardesty, C.A. Hostetler, J.L. Jimenez, A.O. Langford, 825 E. McCauley, S.A. McKeen, L.T. Molina, A. Nenes, S.J. Oltmans, D.D. Parrish, J.R. 826 Pederson, R.B. Pierce, K.A. Prather, P.K. Quinn, J.H. Seinfeld, C.J. Senff, A. Sorooshian, 827 J. Stutz, J.D. Surratt, M. Trainer, R. Volkamer, E.J. Williams, and S.C. Wofsy (2013), 828 829 The 2010 California Research at the Nexus of Air Quality and Climate Change (CalNex) field study. J. Geophys. Res., In Press. 830
- Schade, G.W., and P.J. Crutzen (1995), Emission of aliphatic amines from animal husbandry and
  their reactions: Potential source of N<sub>2</sub>O and HCN, *J. Atmos. Chem.*, 22 (3), 319-346.
- Setyan, A., Q. Zhang, M. Merkel, W.B. Knighton, Y. Sun, C. Song, J.E. Shilling, T.B. Onasch,
  S.C. Herndon, D. Worsnop, J.D. Fast, R.A. Zaveri, L.K. Berg, A. Wiedensohler, B.A.
  Flowers, M.K. Dubey, and R. Subramanian (2012), Characterization of submicron
  particles influenced by mixed biogenic and anthropogenic using high-resolution aerosol
  mass spectrometry: Results from CARES, *Atmos. Chem. Phys.*, *12*, 8131-8156.
- Shields, L.G., D.T. Suess, and K.A. Prather (2007), Determination of single particle mass
  spectral signatures from heavy-duty diesel vehicle emissions for PM<sub>2.5</sub> source
  apportionment, *Atmos. Environ.*, 41 (18), 3841-3852.
- Silva, P.J., D.Y. Liu, C.A. Noble, and K.A. Prather (1999), Size and chemical characterization of
   individual particles resulting from biomass burning of local Southern California species,
   *Environ. Sci. Tech.*, 33 (18), 3068-3076.
- Silva, P.J., and K.A. Prather (2000), Interpretation of mass spectra from organic compounds in aerosol time-of-flight mass spectrometry, *Anal. Chem.*, 72 (15), 3553-3562.
- Smith, J.N., M.J. Dunn, T.M. VanReken, K. Iida, M.R. Stolzenburg, P.H. McMurry, and L.G.
   Huey (2008), Chemical composition of atmospheric nanoparticles formed from

- nucleation in Tecamac, Mexico: Evidence for an important role for organic species in nanoparticle growth, *Geophys. Res. Lett.*, *35* (4), L04808, doi:10.1029/2007GL032523.
- Sodeman, D.A., S.M. Toner, and K.A. Prather (2005), Determination of single particle mass
  spectral signatures from light-duty vehicle emissions, *Environ. Sci. Tech.*, *39* (12), 45694580.
- Song, X.H., P.K. Hopke, D.P. Fergenson, and K.A. Prather (1999), Classification of single
   particles analyzed by ATOFMS using an artificial neural network, ART-2a, *Anal. Chem.*,
   71 (4), 860-865.
- Sorooshian, A., S.N. Murphy, S. Hersey, H. Gates, L.T. Padro, A. Nenes, F.J. Brechtel, H.
  Jonsson, R.C. Flagan, and J.H. Seinfeld (2008), Comprehensive airborne characterization
  of aerosol from a major bovine source, *Atmos. Chem. Phys.*, 8 (17), 5489-5520.
- Sorooshian, A., L.T. Padro, A. Nenes, G. Feingold, A. McComiskey, S.P. Hersey, H. Gates,
  H.H. Jonsson, S.D. Miller, G.L. Stephens, R.C. Flagan, and J.H. Seinfeld (2009), On the
  link between ocean biota emissions, aerosol, and maritime clouds: Airborne, ground, and
  satellite measurements off the coast of California, *Glob. Biogeochem. Cy.*, 23, GB4007,
  doi:10.1029/2009GB003464.
- Spencer, M.T., L.G. Shields, D.A. Sodeman, S.M. Toner, and K.A. Prather (2006), Comparison
   of oil and fuel particle chemical signatures with particle emissions from heavy and light
   duty vehicles, *Atmos. Environ.*, 40 (27), 5224-5235.
- Srivastava, A., M.E. Pitesky, P.T. Steele, H.J. Tobias, D.P. Fergenson, J.M. Horn, S.C. Russell,
  G.A. Czerwieniec, C.S. Lebrilla, E.E. Gard, and M. Frank (2005), Comprehensive
  assignment of mass spectral signatures from individual Bacillus atrophaeus spores in
  matrix-free laser desorption/ionization bioaerosol mass spectrometry, *Anal. Chem.*, 77
  (10), 3315-3323.
- Sullivan, R.C., and K.A. Prather (2005), Recent advances in our understanding of atmospheric
   chemistry and climate made possible by on-line aerosol analysis instrumentation, *Anal. Chem.*, 77 (12), 3861-3885.
- Toner, S.M., L.G. Shields, D.A. Sodeman, and K.A. Prather (2008), Using mass spectral source
  signatures to apportion exhaust particles from gasoline and diesel powered vehicles in a
  freeway study using UF-ATOFMS, *Atmos. Environ.*, 42 (3), 568-581.
- Twohy, C.H., M.D. Petters, J.R. Snider, B. Stevens, W. Tahnk, M. Wetzel, L. Russell, and F.
  Burnet (2005), Evaluation of the aerosol indirect effect in marine stratocumulus clouds:
  Droplet number, size, liquid water path, and radiative impact, *J. Geophys. Res.-[Atmos.]*, *110* (D8), doi:10.1029/2004JD005116.
- Vogt, R., P.J. Crutzen, and R. Sander (1996), A mechanism for halogen release from sea-salt aerosol in the remote marine boundary layer, *Nature*, *383* (6598), 327-330.
- Weiss-Penzias, P., E. Williams, B. Lerner, T. Bates, C. Gaston, K. Prather, A. Vlasenko, and S.M. Li (2013), Shipboard Measurements of Gaseous Elemental Mercury along the Coast
  of Central and Southern California, *Journal of Geophysical Research-[Atmos]*, *118*, DOI:
  10.1029/2012JD018463.
- Wenzel, R.J., D.Y. Liu, E.S. Edgerton, and K.A. Prather (2003), Aerosol time-of-flight mass
   spectrometry during the Atlanta Supersite Experiment: 2. Scaling procedures, *J. Geophys. Res.-[Atmos.]*, 108 (D7), 8427,doi:10.1029/2001JD001563.

- Whiteaker, J.R., D.T. Suess, and K.A. Prather (2002), Effects of meteorological conditions on
  aerosol composition and mixing state in Bakersfield, CA, *Environ. Sci. Tech.*, 36 (11),
  2345-2353.
- Zhang, Q., C.O. Stanier, M.R. Canagaratna, J.T. Jayne, D.R. Worsnop, S.N. Pandis, and J.L.
  Jimenez (2004), Insights into the chemistry of new particle formation and growth events
  in Pittsburgh based on aerosol mass spectrometry, *Environ. Sci. Tech.*, 38 (18), 47974809.
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900 Table 1: Meteorological and Gas Phase Data for 6 Different Air Mass Transport 901 Conditions. Day of Year (DOY), date and time (UTC), latitudinal and longitudinal range, 902 average meteorological conditions (wind speed, RH, air temperature), and average gas-phase 903 concentrations (radon, SO<sub>2</sub>, ozone, and the toluene/benzene ratio) for measurements made during 904 each of the 6 different time periods are highlighted.

Period Location/ Air Mass	ion/ Iass	Day of Year (DOY)	Date and Time (UTC)	Latitude	Latitude Longitude Speed (m/s)	Wind Speed (m/s)	RH (%)	Air Temp (°C)	Radon SO <sub>2</sub> Ozone Toluene/ (mBq/m <sup>3</sup> ) (ppbv) (ppbv) Benzene	SO <sub>2</sub> (ppbv)	Ozone (ppbv)	Ozone Toluene/ (ppbv) Benzene
Rive	Riverside	135.5- 137.375	5/15/2010 12:00- 5/17/2010 9:00	33.14N to 34.02N	118.33W to 119.22W	2.93	89	13.2	2500	0.07	47.1	N/A
Staş P	Stagnant/ Ports	137.375- 140.5	5/17/2010 9:00- 5/20/2010 12:00	33.67N to 34.38N	118.22W to 119.69W	3.86	91	13.4	1060	0.06	39.0	0.40
C M	Marine/ Coastal	142-146	5/22/2010 0:00- 5/26/2010 0:00	33.31N to 33.95N	118.07W to 118.93W	5.18	72	13.9	1550	0.14	35.1	0.85
Po L/	Ports of LA/LB	146.33- 147.875	5/26/2010 8:00- 5/27/2010 21:00	33.53N to 33.77N	33.53N to 118.1W to 33.77N 118.5W	2.40	73	16.0	740	3.40	29.3	1.89
In V	Inland/ Valley	149- 151.875	5/29/2010 0:00- 5/31/2010 21:00	33.5N to 34.4N	118.17W to 119.85W	3.29	83	15.2	3200	0.45	43.6	0.84
Bay	Bay Area/ sacramento	Bay Area/ 154.625- Sacramento 157.625	6/3/2010 15:00- 6/6/2010 15:00	38.02N to 38.56N	121.55 W to 122.16W	4.53	69	20.9	1550	0.42	22.8	1.74

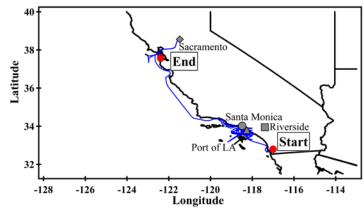
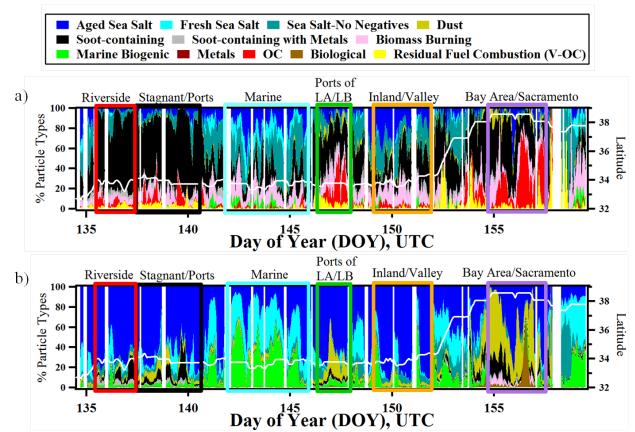


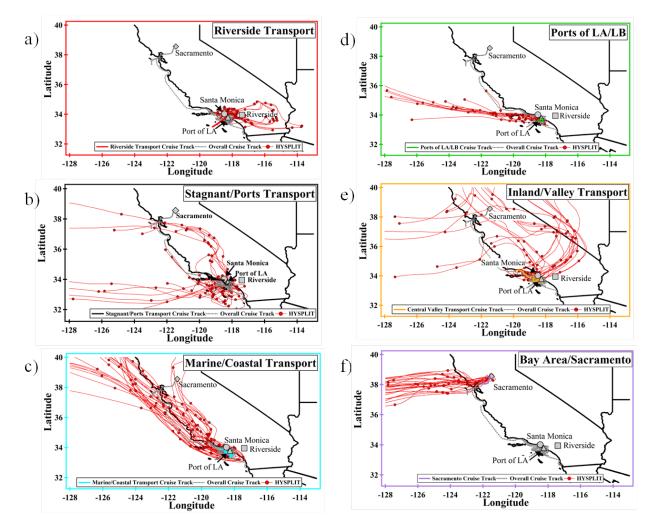
Figure 1: Cruise track for the R/V Atlantis (blue line) along the California coast during CalNex.
The Port of Los Angeles (grey triangle), the Santa Monica area (grey dot), Riverside (grey
square) and the Sacramento area (grey diamond) are shown along with the start and end points of
the cruise.





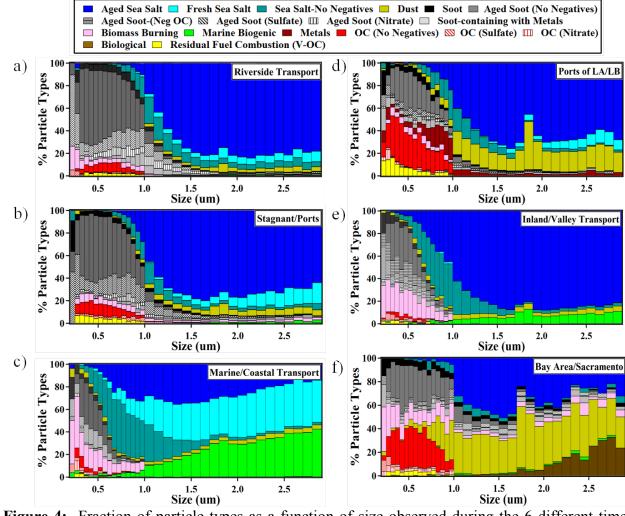
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**Figure 2:** Hourly temporal profile of single-particle mixing-state observed by ATOFMS in UTC as a function of day of year (DOY) and latitude (white line). The (a) top and (b) bottom panels show the single-particle chemistry for submicron particles (0.2-1.0  $\mu$ m) and supermicron particles (1.0-3.0  $\mu$ m), respectively. Colored boxes highlight 6 different periods when differences in particle composition were observed due to different meteorological conditions, gas-phase concentrations, and aging processes.

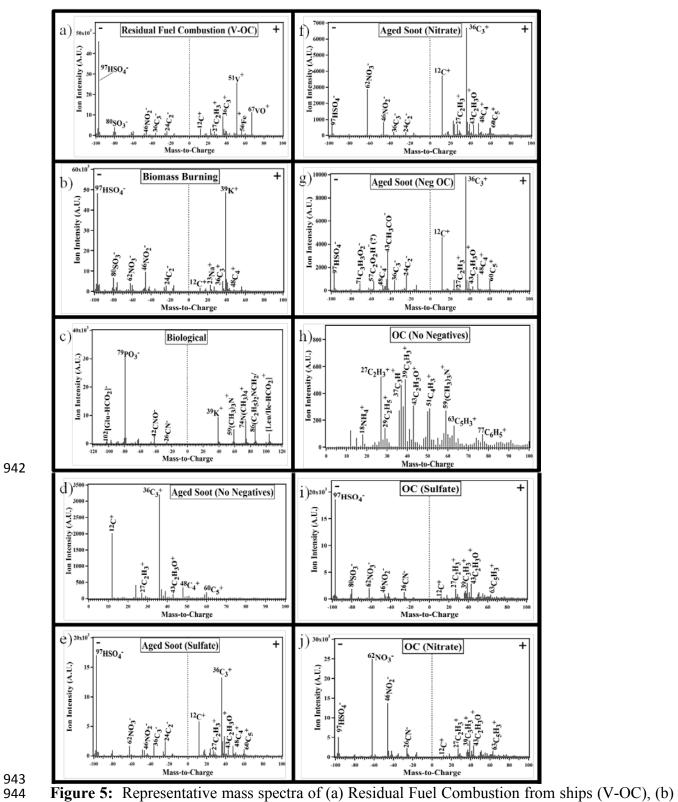


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Figure 3: 48-hour HYSPLIT air mass back-trajectories at 500 m (red lines) shown during the 930 cruise (grey dotted line) corresponding to (a) Period 1: Riverside Transport boxed in red, (b) 931 Period 2: Stagnant/Ports Transport boxed in black, (c) Period 3: Marine/Coastal Transport 932 boxed in cyan, (d) Period 4: Ports of LA/LB boxed in green, (e) Period 5: Inland/Valley 933 Transport boxed in orange, and (f) Period 6: Bay Area/Sacramento boxed in purple. Red dots 934 on the HYSPLIT trajectories denote 12 hour increments. The Port of Los Angeles (grey 935 triangle), the Santa Monica area (grey dot), Riverside (grey square) and the Sacramento area 936 (grey diamond) are also shown. 937



938Size (um)Size (um)939Figure 4: Fraction of particle types as a function of size observed during the 6 different time940periods. Submicron particles (0.2-1.0 μm) are plotted in 0.05 μm bins while supermicron941particles (1.0-3.0 μm) are plotted in 0.1 μm bins.



Biomass Burning, (c) Biological, (d) Aged Soot (No Negatives), (e) Aged Soot (Sulfate), (f)

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946	Aged Soot	(Nitrate), (g	) Aged Soot	(Neg OC),	(h) OC (	No Negatives),	(i) OC	(Sulfate), (j)	) OC

- 947 (Nitrate) particles are shown. For mass spectra containing both positive and negative ions,
- 948 dashed lines separate negative ions (left side) and positive ions (right side).

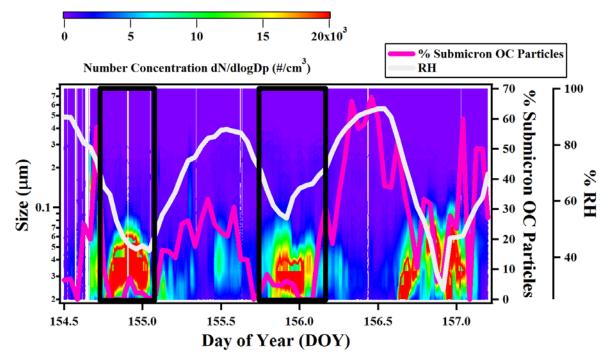


Figure 6: Size distributions of particle number concentrations as a function of size and DOY
are shown on a log scale. The percentage of submicron OC particles detected by ATOFMS
(pink line) and RH (grey line) are shown as a function of DOY. Time periods when new particle
formation events were observed are shown in black boxes.

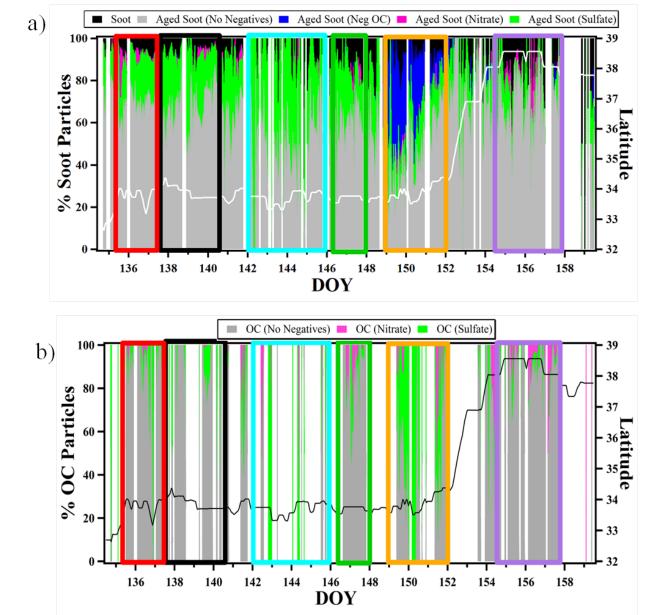
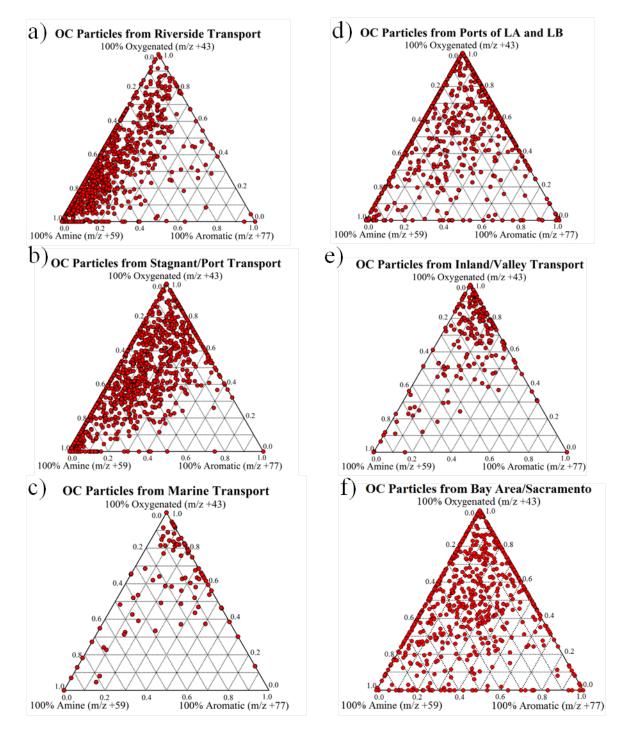


Figure 7: Hourly temporal trends in carbonaceous mixing-state. The top panel (a) shows hourly temporal trends for different soot particle types: Soot (black), Aged Soot (No Negatives) (grey), Aged Soot (Sulfate) (green), Aged Soot (Nitrate) (pink), and Aged Soot (Neg OC) (blue) in addition to latitude (white line). The bottom panel (b) shows hourly temporal trends for different organic particle types: OC (No Negatives) (grey), OC (Sulfate) (green), and OC (Nitrate) (pink) in addition to latitude (black line). The 6 different time periods are boxed.





**Figure 8:** Ternary plots for individual OC particles observed for the 6 different time periods. The top corner of the ternary plots corresponds to OC particles containing only the oxygenated organic peak ( ${}^{43}C_2H_3O^+$ ), the left bottom corner denotes OC particles containing only the amine

970 peak ( $^{59}(CH_3)_3N^+$ ), and the right bottom corner corresponds to OC particles containing only the 971 aromatic peak ( $^{77}C_6H_5^+$ ).